



Integrating palaeoclimate time series with rich metadata for uncertainty modelling: strategy and documentation of the PalMod 130k marine palaeoclimate data synthesis

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Abstract. Palaeoclimate data hold the unique promise of providing a long-term perspective on climate change and as such can serve as an important benchmark for climate models. However, palaeoclimate data have generally been archived with insufficient standardisation and metadata to allow for transparent and consistent uncertainty assessment in an automated way. Thanks to improved computation capacity, transient palaeoclimate simulations are now possible, calling for data products containing multi-parameter time series rather than information on a single parameter for a single time slice. Efforts are underway to simulate a complete glacial–interglacial cycle using general circulation models (<https://www.palmod.de/>, last access: 6 May 2020), and to confront these simulations with palaeoclimate data, we have compiled a multi-parameter marine palaeoclimate data synthesis that contains time series spanning 0 to 130 000 years ago. We present the first version of the data product that focuses exclusively on time series for which a robust chronology based on benthic foraminifera $\delta^{18}\text{O}$ and radiocarbon dating is available. The product contains 896 time series of eight palaeoclimate parameters from 143 individual sites, each associated with rich metadata, age–depth model ensembles, and information to refine and update the chronologies. This version contains 205 time series of benthic foraminifera $\delta^{18}\text{O}$; 169 of benthic foraminifera $\delta^{13}\text{C}$; 131 of seawater temperature; 174 and 119 of planktonic foraminifera $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$; and 44, 38 and 16 of carbonate, organic carbon and biogenic silica content, respectively. The data product is available in three formats (R, LiPD and netCDF) facilitating use across different software and operating systems and can be downloaded at <https://doi.org/10.1594/PANGAEA.908831> (Jonkers et al., 2019). This data descriptor presents our data synthesis strategy and describes the contents and format of the data product in detail. It ends with a set of recommendations for data archiving.

1 Introduction

Global climate has varied dramatically over the last glacial–interglacial cycle. Since the previous interglacial (approximately 130 000 years ago) the Earth had slowly been cooling until the Last Glacial Maximum (LGM; approximately 21 000 years ago). This cooling was associated with the growth of massive ice sheets in North America and Eurasia, leading to a sea level drop of about 120 m (Waelbroeck et al., 2002) and pronounced climate variability on millennial timescales (Voelker and workshop participants, 2002). From the LGM, the Earth warmed rapidly until the onset of the current relatively stable warm period, the Holocene (Shakun et al., 2012). The ultimate cause of the large-scale variations in the Earth's climate is changes in the orbit of the Earth around the Sun (Hays et al., 1976). However, complex feedback and non-linear mechanisms, involving ocean (atmosphere, cryosphere) circulation and biogeochemical cycles, are required to explain how slow changes in the orbital configuration led to the observed evolution of global climate and how these processes led to the manifestation of abrupt climate change.

For these reasons the last glacial–interglacial cycle has been a key target for palaeoclimate modelling. Initially this only involved equilibrium simulations for key time slices, such as the LGM, or transient simulations for short periods, such as the last millennium. The motivation to simulate past climate states is given by the possibility of palaeoclimate data serving as a benchmark for the models. Indeed, this possibility contributed to the development of large palaeodata syntheses (CLIMAP project members, 1981; MARGO project, 2009). The time-slice modelling approach is still being pursued; for example in phase 4 of the Paleoclimate Modelling Intercomparison Project (PMIP), four of the five target intervals fall into the time frame of the last glacial cycle (Kageyama et al., 2018). However, with increasing computing power, the focus is now shifting towards transient climate simulations (Liu et al., 2009; Latif et al., 2016), and the simulation of the last deglaciation is now also considered in the PMIP protocol (Ivanovic et al., 2016).

This development calls for a different type of palaeodata synthesis, with its focus on time series rather than on time slices. Time series of climate data are needed to evaluate aspects of transient simulations that are not available in equilibrium simulations, such as rates of change, phase relationships and spectral properties of climate variability. It is also clear that an evaluation in multi-parameter space using different aspects of the climate system and multiple proxies will be more powerful and diagnostic (Kurahashi-Nakamura et al., 2017), calling for multi-parameter synthesis products.

Observations of the evolution of past climate are based on proxies (measurable approximations of climate-related variables) and hence are, by definition, indirect. Comparison of proxy-based reconstructions with climate model simulations is therefore far from straightforward, as discrepancies may

arise from both model and proxy uncertainty. Proxy uncertainty derives from reconstruction uncertainty (related to calibration, recording bias, archive specifics and instrumental approach) and chronological uncertainty. The latter is particularly relevant to the comparison of transient climate change, and chronological uncertainty thus requires a comprehensive treatment in data syntheses of palaeoclimate time series.

Accounting for proxy uncertainties in a comprehensive and transparent manner requires not only expert knowledge but also the availability of extensive metadata in addition to the proxy data. However, due to a lack of standardisation and inconsistent archiving of metadata, synthesising palaeoclimate data in a way that allows for robust uncertainty assessment remains challenging and time consuming. Efforts are underway to alleviate these challenges. The largest palaeoclimate data repositories (World Data Service for Paleoclimatology, operated by the national centres for environmental information (NCEIs) at NOAA, and PAN-GAEA) are both striving for more standardisation and to store data in (more) machine-readable formats. In addition, standardisation is progressing through the use of existing data formats from other communities (netCDF; Langner and Mulitza, 2019) as well as the implementation of new data formats specifically targeted to palaeodata (Linked Paleo Data (LiPD); McKay and Emile-Geay, 2016). At the same time there is ongoing discussion on data and metadata requirements and standards (Khider et al., 2019). Traceability of datasets is also improved through data citations, not only ensuring that data producers receive proper credit for their work but also allowing for better linking of different datasets. Nevertheless, these initiatives have only recently been emerging and the majority of the palaeoclimate data remains inconsistently formatted, non-standardised and scattered over various data repositories. The need for synthesis products and documentation of potential synthesis approaches is therefore as large as ever.

Here we present the first version of a new multi-proxy marine palaeoclimate data synthesis that covers the past 130 000 years developed within the German climate modelling initiative PalMod (Latif et al., 2016). We focus on the ocean as it is a large reservoir of heat and CO₂ and allows for global coverage with consistent chronological control. This synthesis goes beyond the time frame of many existing multi-proxy and multi-parameter data syntheses (PAGES2k Consortium et al., 2017; Routson et al., 2019), expands existing data products that provide long palaeoclimate time series to multiple parameters (Shakun et al., 2012; Marcott et al., 2013; Peterson and Lisiecki, 2018; Snyder, 2016), and is based on a strategy of semi-automated data harvesting (Cartapanis et al., 2016). This version of the synthesis contains data on nine climate-sensitive parameters: benthic and planktonic foraminifera stable oxygen and carbon isotopes, seawater temperature, radiocarbon and bulk sediment carbonate, organic carbon, and biogenic silica content.

In this paper we describe our synthesis approach, the contents and structure of version 1.0.0 of the data, plans for future updates, and recommendations for archiving new data and retrieving dark data in a way that allows for optimal future reuse. The data product is intended to be used to investigate spatio-temporal changes in a multi-parameter domain. Thanks to rich metadata that allow for the rigorous quantification of reconstruction uncertainties, we also envision that this data product will provide the building blocks for intelligent palaeoclimate data model comparison (Weitzel et al., 2019), for instance through proxy system modelling (Dolman and Laepple, 2018) or data assimilation (Breitkreuz et al., 2019).

The structure of this data descriptor is as follows. Section 2 describes the synthesis strategy, including the data discovery approach, standardisation and age modelling. In Sect. 3 we provide general information on palaeoclimate proxies from marine-sediment archives that is used to guide the metadata selection. Section 4 details the structure of the database, and the contents of version 1.0.0 are outlined in Sect. 5. The formats of the data product, future plans and versioning are described in Sects. 6 and 7. Section 8 describes where the data can be accessed. In the last section, Sect. 9, we reflect on the data synthesis effort and provide recommendations for data archiving and data rescue.

2 Data synthesis strategy

Our data product focuses on time series from marine-sediment archives. A single marine-sediment archive (sediment core) can be used for measurements of different parameters, each providing information on different aspects of the environmental conditions at the time of deposition. However, for the purpose of analysis, the various proxy time series must refer to a single age–depth model for the sediment core they are derived from. For this reason, the basis of our synthesis is formed by a collection of sediment cores, each associated with its own age–depth model.

Marine sediments are dated using absolute age controls, where specific layers are dated using, for instance, radiocarbon, tephra or palaeomagnetic properties, and/or relative age controls, where time series are aligned based on the hypothesised synchronicity of the changes recorded by some properties of the sediment. A well-established hypothesis-based age modelling approach with a solid theoretical basis is the alignment of benthic foraminifera stable-oxygen-isotope ratio ($\delta^{18}\text{O}$) time series (Lisiecki and Raymo, 2005). We thus base our chronological framework on a combination of radiocarbon dates and benthic foraminifera $\delta^{18}\text{O}$ and have selected time series where both parameters are available as the foundation of this data product. This approach of blending absolute and relative age controls is required to provide age–depth models for sediment cores that extend beyond the radiocarbon dating range ($\sim 40\,000$ years). If available, further

Table 1. Palaeoclimate parameters in the PalMod 130k marine data synthesis.

Parameter
Benthic foraminifera $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$
Planktonic foraminifera $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$
Seawater temperature*
Radiocarbon
Carbonate content
Total organic carbon content
Biogenic silica content

* Inferred from various proxies (foraminifera Mg/Ca, alkenones, microfossil assemblages).

proxy time series were then added, thus ensuring a common chronology among all proxy time series measured on the same sediment core.

We selected palaeoclimate parameters to synthesise the following discussion with climate modellers within the PalMod project. The high-priority selection includes both physically and biogeochemically relevant parameters, of which some are based on measurements that can be compared with climate model output using (forward proxy) models (e.g. benthic $\delta^{18}\text{O}$) and others represent inferred parameters that can be compared with model output more directly but for which proxy models are still in their infancy (e.g. temperature based on foraminifera Mg/Ca). Also considered in parameter selection was the expected spatial and temporal coverage of data availability as well as the existence of previous data products. The high-priority parameters for which data are presented here are listed in Table 1. If available, raw data were synthesised and in cases where raw data were not available and it was possible to derive the raw data from the inferred palaeoclimate data, raw data were back-calculated. Raw data time series obtained in this way are flagged with a note describing the calculation.

We note that our approach of first building the stratigraphic framework based on radiocarbon dates and benthic foraminifera $\delta^{18}\text{O}$ means that the synthesis is not necessarily comprehensive as it does not include time series where one of the parameters of interest has been measured but where the components of the stratigraphic framework are not available. However, at this stage, we opted to include only sediment cores where an age modelling strategy that is consistent and comparable across the entire data product could be achieved.

2.1 Data discovery

In principle, data synthesis can proceed by expansion or reduction (Fig. 1). The first, more traditional, approach relies on expert knowledge of what data are available and/or on asystematic literature search. In this approach the synthesis grows by including more data until sufficient data that meet inclusion criteria are compiled. In this way, a lot of time is

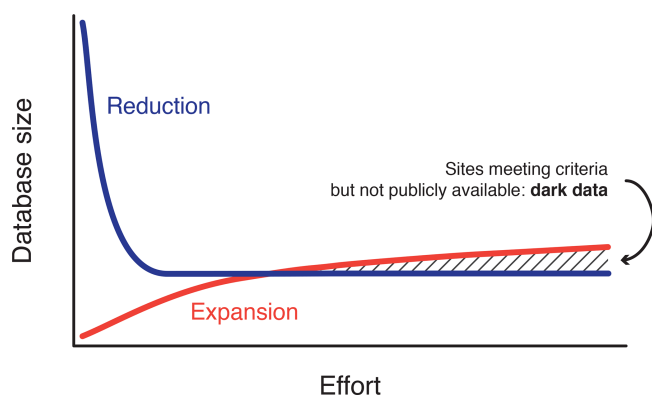


Figure 1. Data synthesis approaches. In the expansion approach the database size increases slowly as records are added. The database size follows an opposite pathway using the reduction approach and reaches a stable size more quickly, with less effort. Since the expansion approach is not restricted to data that are available in the public domain, this approach may lead to a database that includes data that are not publicly available (dark data). The reduction approach on the other hand is arguably more objective, can be automated and is therefore more efficient. This approach also encourages good data stewardship.

spent on discovering and retrieving datasets, and it is possible that valuable, but less exposed, data are missed. On the other hand, this approach has a chance of uncovering dark data that are not publicly available (Fig. 1).

The second approach starts from a large and crude synthesis of data from public sources and proceeds by weeding out data that do not meet criteria for inclusion in the data product. This approach faces different challenges: making sure that the initial bulk database is comprehensive (efficient data mining) and assuring that the data filtering is efficient (fast and accurate). In contrast to the expansion approach, this reduction approach cannot discover dark data. However, it is more objective (less reliant on expert knowledge), can be automated more easily and focuses on data that are already in the public domain so that no time is lost to finding data that ultimately prove unavailable. Because this second approach focuses on data that are publicly available, it also rewards and encourages good data stewardship.

In theory, both approaches can lead to a similarly sized and exhaustive synthesis, but they differ in the allocation of effort (Fig. 1). In practice both approaches are often combined, especially towards the end of a synthesis project, when the data product is benchmarked against existing syntheses.

2.2 Synthesis

2.2.1 Initial synthesis

We followed the reduction approach and used a semi-automated pipeline to compile data from public sources. Keywords (Supplement) were used to make lists of

URLs of potentially relevant data on <https://pangaea.de> (last access: September 2016), and the linked files were then downloaded in bulk ($n = 108\,239$). A slightly different approach was followed for the NCEI archive. Here, all files that were machine-readable at the time of download (September 2016, $n = 1925$) were obtained from the FTP server (ftp://ftp.ncdc.noaa.gov/pub/data/paleo/paleocean/sediment_files/complete, last access: September 2016). Custom scripts in R were used to put all data in a common format and merge time series that could be unambiguously assigned to the same core (based on name and x , y , z position). This resulted in a mixture of records that were merged to the same core and those that could not be, either because there was only one data file for the core or because of ambiguous labelling. We refer to the locations of these records as “sites”. In order to facilitate the analysis (filtering) of the sites, a uniform attribution of the various parameter names had to be developed. Because no standardised names exist for palaeoclimate parameters, the uniform attribution required the development of attribution libraries for each desired palaeoclimate parameter. The initial synthesis contained time series from 38 511 sites.

2.2.2 Data reduction and standardisation of ontologies

The initial synthesis was reduced by removing non-marine sites (using elevation) and further constrained by only considering sites where at least one data point of any of the parameters measured in that core fell within the target time frame (disambiguating age units in the synonym library of the category “age”) and the site had benthic oxygen isotope data. This resulted in 781 sites. At this stage, no criteria for length or resolution were applied but we prioritised processing time series that we estimated to contain at least 50 data points within the 130 000-year timeframe. Further data processing started with dereplication of the selected sites. This was necessary because no standards exist for the naming of cores and the repositories store data with different renditions of the same core name, sometimes even associated with erroneous geographic coordinates. This process was carried out manually and proceeded by constructing a list of disambiguated sites through sequential one-by-one comparison. Where a strict synonym was found (different labels for the same core but the same data), only unique data were retained. At this stage, disambiguation of site names was only performed for sites that had at least benthic $\delta^{18}\text{O}$, so time series of other parameters, which were associated with inconsistent core labels, could have been missed in the synthesis. However, those sites are contained in the initial bulk synthesis and are hence not lost but will be salvaged in updates of the data product (see Sect. 7).

Further steps required a manual standardisation of the names of the parameters and their attributes (such as the species name that was analysed for oxygen isotopes). This was accomplished by deciding on a final, uniform list of pa-

rameters and associated metadata and their possible values. Original parameter names were preserved to allow for cross-checking. By metadata, we refer to aspects of the individual parameters that were deemed essential to facilitate a meaningful analysis in a palaeoclimatic context, considering potential sources of uncertainty, such as species of foraminifera analysed or the calibration equation used for palaeotemperature estimates. The list of metadata is provided in Table 2. The standardisation was accompanied by further dereplication of individual time series that were already associated with the same site name but archived more than once.

2.2.3 Metadata and chronology

Subsequently, as far as possible, metadata values were added manually, often by scraping the information from the original publication. Next, all time series from a single site (core) were put on a common depth scale to allow for age modelling. Data that could not be put on a depth scale were excluded from the synthesis. This was the case where parameter values were archived only against age and where no other data file was available that allowed unambiguous reassignment to depth. Ambiguity also resulted from the use of multiple (composite) depth scales for the same archive. Finally, chronological data (all absolute markers, including radiocarbon dates and associated metadata) were manually added, where necessary also by consulting the original publications. Throughout the process, publication information (digital object identifier (DOI) or, if not available, full bibliographic details) and the data source (URL and/or DOI) were preserved in order to trace the source of the data. This applies both to the source of the individual data files from repositories and to the sources of the metadata and chronological data.

2.3 Age modelling

Whereas the initial steps of data discovery and synthesis could rely on published chronology, analysis of the complete dataset requires the development of a common chronological framework. This framework must be constructed in a way that not only allows for a consistent method of assignment of ages to depths within each core but also allows for the consistent and quantitative assessment of age uncertainty. To this end, we follow an approach that combines absolute ages (radiocarbon ages, tephra layers and palaeomagnetic events) with $\delta^{18}\text{O}$ stratigraphy. As a result, our age models may differ from those reported in the original publication(s). This does not mean that the updated age models are better (constrained), but they are constructed in a way that allows for applicability and consistency across the synthesis. The consistent approach allows for an assessment of age uncertainty jointly for all records by a Bayesian approach, generating ensembles of sedimentation histories consistent with the available age control points for each core, allowing for uncer-

tainty estimates at each depth by considering the distribution of ages given by the ensemble.

With respect to the reporting of the chronology, we follow a transparent approach, preserving the initial age model and providing the new age models as well as all information needed to revise or update the new age models. In the final step, the age information from absolute ages and $\delta^{18}\text{O}$ tie points was combined and the age model and its uncertainty was assessed in a Bayesian framework using “Bacon” (Blaauw and Christen, 2011). The entire age modelling routine was carried out in PaleoDataView (PDV; Langner and Mulitza, 2019).

To ensure a common chronological framework for all time series in the synthesis, radiocarbon ages were recalibrated using the IntCal13 curve (Reimer et al., 2013). Since reservoir ages vary in space as well as in time, we used reservoir age estimates based on a comprehensive ocean general circulation model (Butzin et al., 2017) to account for this variability in a physically plausible way. To derive the reservoir age and uncertainty for a measured radiocarbon age, PDV (i) extracts all modelled radiocarbon ages from the nearest grid cell in the modelled dataset, (ii) finds all modelled radiocarbon ages that are possible within the error of the measured radiocarbon age, and (iii) takes the mean and the standard deviation of all corresponding reservoir ages to correct for the measured radiocarbon age. By definition, this approach cannot account for processes affecting the reservoir ages on subgrid spatial scales. Given the relatively coarse resolution of the model, this means that processes such as upwelling are not fully accounted for. In addition, no modelled reservoir age data are available for the Mediterranean and Red seas; we use the reservoir ages reported by the authors of the original publication and an assumed uncertainty of 100 years for these basins (five sites). Absolute ages based on North Atlantic tephra layers and palaeomagnetic events were updated and harmonised using Svensson et al. (2008).

In addition, and beyond the ^{14}C dating realm ($\sim 40\,000$ years), the age models rely on manual tuning of the benthic foraminifera $\delta^{18}\text{O}$ time series from each core to regional benthic foraminifera $\delta^{18}\text{O}$ stacks (Lisiecki and Stern, 2016). Stable-isotope stratigraphy in theory provides a range of events to correlate; however, in order to not inflate confidence in the tuned age models and to ensure comparability between different cores, in our approach, the tuning was carried out as far as possible by only matching the position of marine isotope stage boundaries. We updated the age–depth models only for the 0–130 000 years time frame of this synthesis, but data and original age models extending beyond 130 000 years are preserved in the data product. To obtain uncertainty for the age control points obtained by $\delta^{18}\text{O}$ tuning, we used the chronological uncertainty in the $\delta^{18}\text{O}$ stacks, as reported by Lisiecki and Stern (2016). Additional uncertainty associated with the identification of the control points in the individual records or with the assumption on synchronicity

Table 2. Metadata terms.

Name	Description
ParameterOriginal*	Original parameter name as in data repository
Parameter*	Standardised parameter name (see Table 4)
ParameterType*	Parameter type, measured or inferred
ParameterUnit*	
ParameterAnalyticalError	Error based on repeat measurements of standards
ParameterReproducibility	Error based on repeat measurements of samples
Instrument	
Laboratory	
SampleThickness_cm	
Material*	Measurement material or parameter on which inferred parameter is based
Species	
Nshells	
SizeFraction_microm	
Notes	
RecordingSeason	
RecordingDepth	
EquilibriumOffset	
CalibrationEquation	
CalibrationUncertainty	
CalibrationDOI	
TransferFunctionTrainingSet	
TransferFunctionUncertainty	
TransferFunctionDOI	
PublicationDOI*	
Authors	
PublicationTitle	
Journal	
Year	
Volume	
Issue	
Pages	
ReportNumber	
DataDOI	
DataLink*	
RetrievalNumber	For internal use only

* Essential terms.

was ignored, as these are difficult, if not impossible, to quantify.

3 Notes on palaeoclimate proxies in marine-sediment archives and metadata

This synthesis contains climate-sensitive proxy data based on measurements using various biological sensors. It is not the intention here to provide a full overview of marine palaeoclimate proxies and their uncertainties (for this, see for example Hillaire-Marcel and De Vernal, 2007; Moffa-Sánchez et al., 2019), but the fact that the proxies are based on biological sensors means that they are affected by different ecological bias in addition to observational noise. Basic knowledge of the recording system is therefore essential for the interpretation of the data and may aid in explaining differences be-

tween proxies for the same climate parameter. These considerations were also essential to choosing the range of metadata to be recorded alongside each palaeoclimatic parameter to allow for a proxy-specific assessment of uncertainty.

Foraminifera are among the most widely used proxy sensors in palaeoceanography. They are unicellular marine zooplankton. The species used here all build a calcite skeleton that is preserved in the sediment. Foraminifera can be divided into two main groups: benthic foraminifera living at the seafloor level or at shallow depth in the sediment and planktonic foraminifera living in the upper hundreds of metres of the ocean. In the data product, proxies measured on these two groups are clearly distinguished by a benthic or planktonic prefix. The chemical composition of foraminifera reflects environmental conditions of the seawater that the organisms calcified in. For the purpose of this data product,

the parameters of interest are stable-oxygen-isotope, stable-carbon-isotope and Mg/Ca ratios. Stable-oxygen-isotope ratios in foraminifera calcite reflect a combination of temperature and $\delta^{18}\text{O}$ of seawater (Urey, 1948), which is in turn related to ice volume and salinity. Species-specific calibrations exist to quantitatively link $\delta^{18}\text{O}_{\text{foraminifera}}$ and $\delta^{18}\text{O}_{\text{seawater}}$ to temperature (e.g. Marchitto et al., 2014; Bemis et al., 1998). Stable-carbon-isotope ratios ($\delta^{13}\text{C}$) reflect the $\delta^{13}\text{C}$ of the dissolved inorganic carbon in seawater. In particular the benthic foraminifera species *Cibicidoides wuellerstorfi* generally incorporates $\delta^{13}\text{C}_{\text{DIC}}$ without a biological offset and can serve as a tracer of bottom-water $\delta^{13}\text{C}_{\text{DIC}}$ which is commonly used as non-passive circulation tracer (Curry and Oppo, 2005). The $\delta^{13}\text{C}$ of other benthic foraminifera species is generally not indicative of bottom-water $\delta^{13}\text{C}_{\text{DIC}}$, and the $\delta^{13}\text{C}$ of planktonic foraminifera is also influenced by temperature and carbonate ion concentration, rendering interpretation complicated (Spero et al., 1997).

The Mg/Ca ratio in foraminifera calcite can be used to infer calcification temperature and, in combination with $\delta^{18}\text{O}_{\text{foraminifera}}$, the $\delta^{18}\text{O}_{\text{seawater}}$ (Elderfield and Ganssen, 2000). Similar to stable oxygen isotopes, species-specific calibrations exist to quantitatively reconstruct past temperature from Mg/Ca ratios (Anand et al., 2003; Lear et al., 2002), and whenever indicated in the original publication, the calibration is included in the metadata. Carbonate system parameters and salinity have a secondary influence on Mg/Ca ratios in foraminifera calcite (Gray et al., 2018b). Whereas benthic foraminifera live in a generally stable environment, the near-sea-surface habitat of planktonic foraminifera shows large seasonal and vertical gradients. Species-specific seasonal and/or depth habitat preferences may therefore leave a considerable imprint on the proxy signal contained in their shells (Jonkers and Kučera, 2017; Mix, 1987). For all proxies based on foraminifera, it is relevant to record the species as well as the number of individuals that were pooled for geochemical analysis. The latter is because the short lifespan and variable habitat of foraminifera species cause large variability among individuals. Planktonic foraminifera shell size may for several reasons also affect their chemistry (Jonkers et al., 2013; Friedrich et al., 2012) as well as their assemblage composition (Al-Sabouni et al., 2007). Therefore, the size fraction of the analysed shells was included in the metadata whenever this information was available.

Besides planktonic foraminifera Mg/Ca ratios, the $\text{U}^{\text{K}'}37$ unsaturation index can provide information about near-sea-surface temperature. The $\text{U}^{\text{K}'}37$ index is based on the relative degree of unsaturation of C_{37} alkenones, which is linearly related to temperature (Prahl et al., 1988). Alkenones are produced by coccolithophores, marine phytoplankton living in the photic zone. The production of alkenones is in many regions not constant during the year, thus potentially causing a seasonal recording bias in the $\text{U}^{\text{K}'}37$ temperature proxy (Rosell-Melé and Prahl, 2013). Several calibrations exist that relate the index to sea surface temperature, and if the calibra-

tion was mentioned in the original publication, it was preserved in the metadata.

A large proportion of the temperature estimates in this data product are based on microfossil (planktonic foraminifera, diatoms, Radiolaria, dinoflagellate cysts) assemblages. These reconstructions are based on a statistical relationship between species assemblages and temperature (Imbrie and Kipp, 1971). In theory, microfossil assemblages can be used to reconstruct temperatures of different seasons or different environmental parameters from the same assemblage. However, it is not always clear that such reconstructions are truly independent (Telford and Birks, 2011). Several different methods exist to relate fossil assemblages to temperature, and researchers often apply more than a single method in their reconstructions to increase confidence (Kucera et al., 2005). When available, these different reconstructions are included in the data product.

The bulk sediment data (CaCO_3 , TOC and BSi) form a category of their own. They are not proxies in the strict sense but properties of the sediment that reflect a combination of export productivity, sedimentation and preservation. However, they can provide crucial information about the ocean–climate system, in particular about biogeochemical cycles (Cartapanis et al., 2016). With the advent of explicit sediment modules in climate models (Heinze et al., 1999; Kurahashi-Nakamura et al., 2020), sediment composition can also be directly compared with model output and potentially provide additional constraints on the simulations.

4 Structure of the database

Following the data synthesis strategy outlined above, we generated a first data product for time series of eight parameters in sediment cores with radiocarbon and benthic $\delta^{18}\text{O}$ stratigraphy. Following the logic of our approach, the synthesis is organised by the physical object from which the records were extracted (cores), here called site, to account for the inclusion of records from spliced cores.

Each site in the data product has information on seven different themes (Fig. 2):

1. Geographic data contain the site name, latitude and longitude (in decimal degrees N and E), elevation or water depth (in metres), and possible notes that are relevant to the site or core as a whole. All fields except notes are essential and always included.
2. Metadata include the original parameter name as given in the online data file, a standardised parameter name, parameter type (measured or inferred), unit, an estimate of analytical error as determined from repeat measurements of a standard and an estimate of reproducibility as determined by repeat measurements on samples. For measured parameters, information on the instrument and laboratory is given. All metadata terms are

listed in Table 2, and an overview of the standardised parameter names is provided in Table 3.

- Chronology data contain raw data on absolute age control points used for age modelling. This includes not only depth, radiocarbon ages including their uncertainty, dated material and laboratory codes but also calendar ages of tephra layers and palaeomagnetic events. Age control points not used in the age model by the authors of the original publication(s) are indicated, and if available, the source (DOI or URL) of the data is shown in addition to the original publication DOI. A complete list of chronology data terms is given in Table 4.
- The actual time series data are provided on a common depth scale. Original age models are preserved alongside the data time series as they may differ for different time series from the same site. The data also contain information on the sample number or label (mainly for DSDP, ODP and IODP cores) for spliced records and sample-specific notes.
- The revised age model contains ages for depths bracketed by age control points (absolute and relative). Mean and median ages are given as well as an uncertainty range (2.5th and 97.5th percentiles) based on the full suite of age model ensembles. No attempts were made to extrapolate the age models beyond the tie points in the 130 000-year time frame, so original age models may extend to either side.
- The Bacon data contain all information to reproduce the revised age model. Besides the ^{14}C and absolute age control points, this includes the tie points for the alignment, the alignment target and all parameters used to construct the age–depth model.
- Age ensembles are provided for further assessment of chronological uncertainty. In order to keep file sizes manageable, 1000 randomly selected age model ensembles are preserved.

5 PalMod 130k marine palaeoclimate data synthesis v1.0.0 contents

The data product contains 896 time series of the palaeoclimate parameters listed in Table 1 from 143 sites. An overview of all datasets used in this synthesis, including URLs to the data, can be found at <https://doi.org/10.5281/zenodo.3739019> (Jonkers et al., 2020). Data sources for each site are listed in the Appendix. By design all sites have both benthic stable-oxygen-isotope and radiocarbon data. The majority of the sites are close to the continents and in the Northern Hemisphere, with a concentration in the North Atlantic Ocean (Fig. 3). This reflects

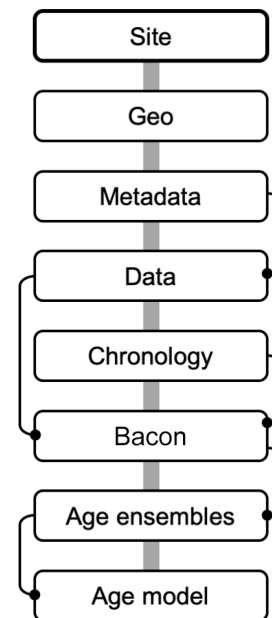


Figure 2. Structure of the PalMod 130k marine palaeoclimate data synthesis. Each file in the database contains information on seven different themes on a single site (sediment core). Links between different themes are indicated.

both research attention as well as the challenges of obtaining sediment cores with high accumulation rates and well-preserved foraminifera. The effect of research focus is also visible in the temporal coverage of the time series, where every parameter is characterised by a clear maximum around the last glacial termination ca. 15 ka (Fig. 4). The median resolution of the time series varies by 2 orders of magnitude but is generally better than one sample per 1000 years and fairly similar among the different parameters (Fig. 5). The updated age models are based on chronological control points (tie points) from radiocarbon dating and absolutely dated layers (using tephra and/or palaeomagnetic event stratigraphy) as well as alignment to the regional benthic $\delta^{18}\text{O}$ stacks. The majority of the time series has a chronological control point at least every 5000 years (Fig. 6). Taken together, the coverage in space, time and across parameters indicates that the PalMod marine palaeoclimate data product allows for analysis of palaeoclimate on a supra-regional scale over the entire 130 000-year time frame.

This data product builds upon previous syntheses. Virtually all of the sites are also part of the benthic foraminifera $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ compilations of Lisiecki and Stern (2016) and Peterson and Lisiecki (2018). Our synthesis, however, also includes data on other palaeoclimate parameters and contains more metadata and information on the age–depth models. Some of the planktonic foraminifera $\delta^{18}\text{O}$ and Mg/Ca time series in the PalMod 130k data product are also included in the Iso2k synthesis effort (Konecky et al., 2018), and a num-

Table 3. Standardised parameter names.

Name	Description
benthic.d18O	Benthic foraminifera $\delta^{18}\text{O}$
benthic.d13C	Benthic foraminifera $\delta^{13}\text{C}$
planktonic.d18O	Planktonic foraminifera $\delta^{18}\text{O}$
planktonic.d13C	Planktonic foraminifera $\delta^{13}\text{C}$
surface.temp	Inferred (near-)sea-surface temperature (based on microfossils, planktonic foraminifera Mg/Ca, $\text{U}^{\text{K}'}37$)
deep.temp	Inferred bottom-water temperature (based on benthic foraminifera Mg/Ca)
CaCO3	Calcium carbonate content
TOC	Total organic carbon content
BSi	Biogenic silica content
DBD	Dry bulk density
IRD	Ice-rafted detritus
planktonic.MgCa	Planktonic foraminifera Mg/Ca ratio
benthic.MgCa	Benthic foraminifera Mg/Ca ratio
UK37	$\text{U}^{\text{K}'}37$ ratio (rare cases where this is not $\text{U}^{\text{K}'}37$ mentioned in notes)
C37.concentration	Alkenone concentration

Table 4. Chronology terms.

Name	Description
ChronType*	Type of absolute chronology tie point (^{14}C , tephra, palaeomag)
ChronDepthTop_cm	
ChronDepthBottom_cm	
ChronDepthMid_cm*	
ChronSampleThickness_cm	
ChronAge_kaBP*	Age of non- ^{14}C tie point (tephra, palaeomag)
ChronAgeError_ka*	
ChronDatedMaterial	Calibrated ^{14}C age
ChronDatedSpecies	
ChronNshellsDated	
Chron14CLabcode	
ChronAge14C_kaBP*	
ChronAge14CError_ka*	
ChronAge14CErrorUp_ka	
ChronAge14CErrorDown_ka	
ChronReservoirAge_ka*	
ChronReservoirAgeError_ka*	
ChronCalibCurve	
ChronCalibAge14C_kaBP	
ChronCalibAge14C1sigLo_ka	
ChronCalibAge14C1sigUp_ka	
ChronAgemodelMethod	
ChronAgeRejected	
ChronNotes	
ChronSource	
ChronDOI*	

* Essential terms.

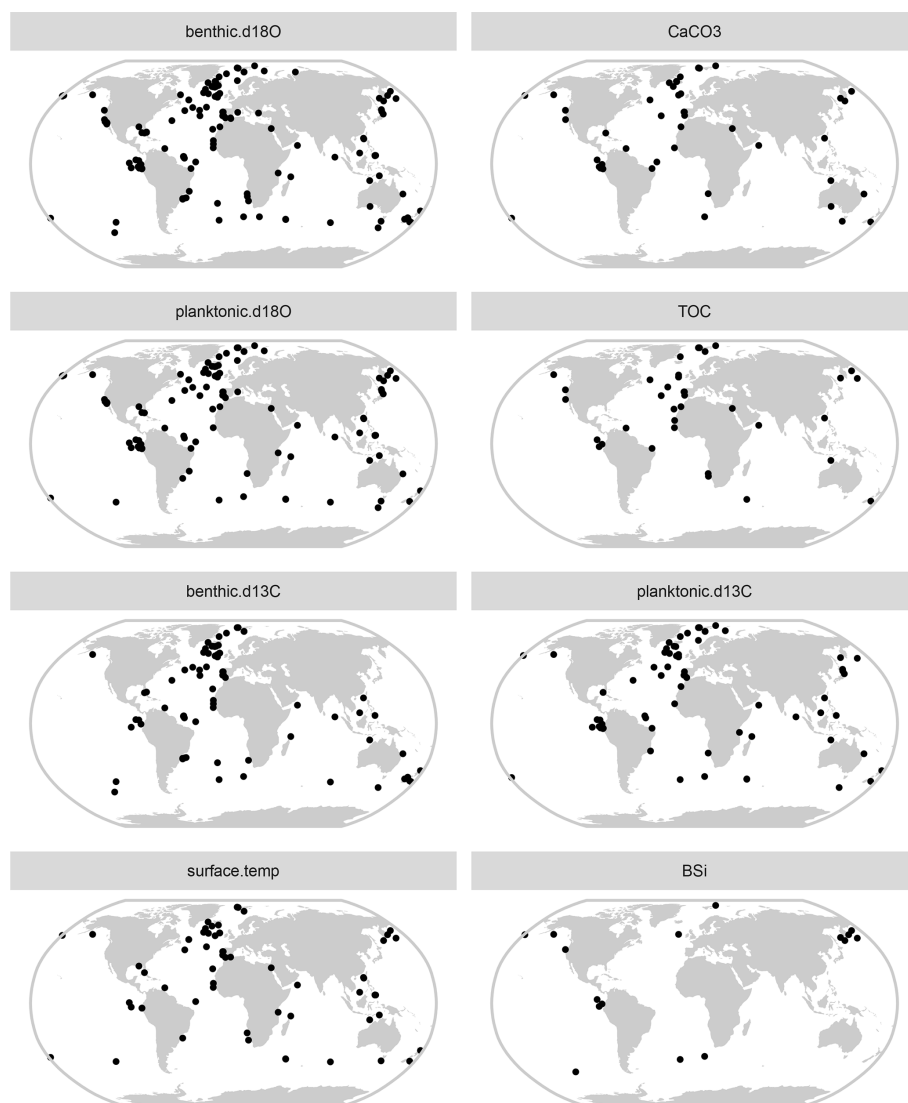


Figure 3. Spatial distribution of sites in version 1.0.0 of the PalMod 130k marine palaeoclimate data synthesis. The distribution of the sites reflects research effort and the possibility of obtaining sediment cores containing well-preserved foraminifera and is hence skewed towards the Northern Hemisphere (Atlantic) and the continental margins. Since no explicit search was carried out for parameters other than benthic foraminifera $\delta^{18}\text{O}$, the distribution of the other parameters is restricted to sites that have benthic foraminifera $\delta^{18}\text{O}$. For benthic foraminifera $\delta^{13}\text{C}$, only sites with data based on the genus *Cibicidoides* are shown.

ber of sea surface temperature time series are also part of the Temperature12k synthesis (Kaufman et al., 2020).

6 Data formats

We provide the data products in three different formats in order to facilitate access and analysis using different software and across operating systems. Given the structure of the formats, each representation is slightly different in its level of metadata detail and the way metadata and data are stored. The differences are described below.

- Since the data product was built using R, the data product is presented in R-readable RDS files that contain for each site a list with data for each theme (Sect. 4). This is the format that is most complete, yet in the interest of memory space it preserves a random selection of 1000 age models from the larger ensemble produced using Bacon. In this format, all data and metadata for each site are contained within a single file. Example scripts (<https://github.com/lukasjonkers/PALMODutils>, last access: 6 May 2020) allow the user to extract a quick overview of the contents of the data product with information on the temporal range, resolution and age control of the time series. Additional code is available

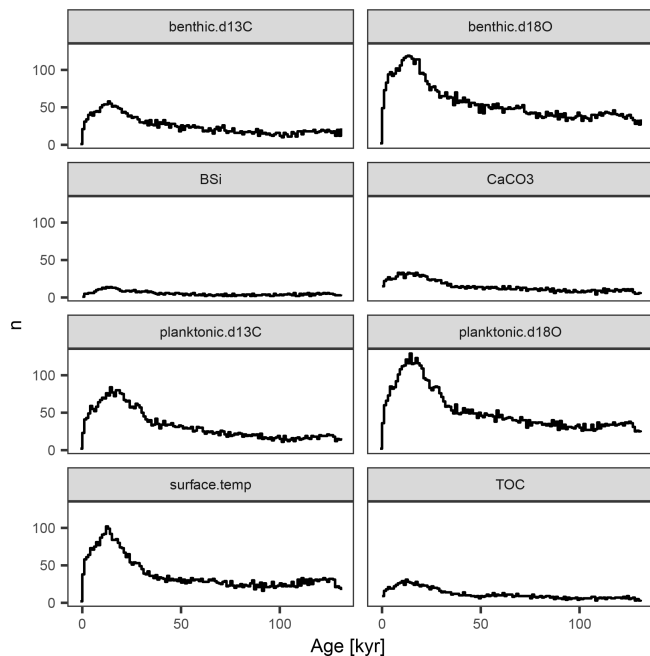


Figure 4. Temporal distribution of time series in version 1.0.0 of the PalMod 130k marine palaeoclimate data synthesis. The number of time series (y axis) is counted per 1000-year bin. The temporal availability of all parameters shows a clear maximum around 15 ka, reflecting the research focus on the last glacial termination. For benthic foraminifera $\delta^{13}\text{C}$, only time series with data based on the genus *Cibicoides* are shown.

to query the data product by parameter, parameter detail, sensor species, temporal range, resolution and age control.

- The data product is also provided in the LiPD format, which is built around JSON-LD and CSV formats and is widely readable across different platforms. As with RDS, all data for each site are presented in a single file. Utilities to interact with LiPD files in R, Matlab and Python are available at <https://github.com/nickmckay/LiPD-utilities> (last access: 6 May 2020).
- Finally, the data are also provided in netCDF format in a way that allows reading with PDV. This means that a single site has separate files for each individual palaeoclimate parameter as well as for the age model. In this format, some metadata are stored as concatenated strings rather than easily searchable attributes. The netCDF format allows however for the storage of the full suite of age model ensembles without excessive file sizes. The PDV software to read and process the data can be downloaded at <https://www.marum.de/en/Stefan-Mulitza/PaleoDataView.html> (last access: 6 May 2020).

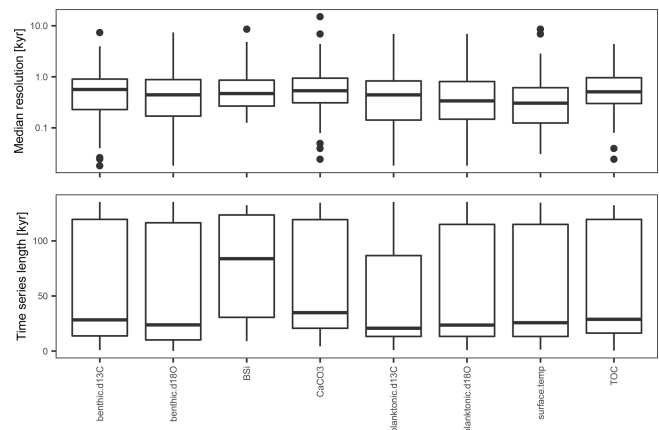


Figure 5. Median resolution of the time series in version 1.0.0 of the PalMod 130k marine palaeoclimate data synthesis. Box-and-whisker plots show the spread of resolution per parameter. For benthic foraminifera $\delta^{13}\text{C}$ the resolution data are restricted to time series containing data measured on shells of the genus *Cibicoides*. For all parameters the median resolution is more than one data point per 1000 years.

7 Future plans and versioning

To increase the spatio-temporal coverage over the entire 130 000-year time frame of the database, updates of this data product will first aim for quantitative growth of the database by adding more time series with chronological control based on benthic foraminifera $\delta^{18}\text{O}$ and absolute age control points other than ^{14}C . If available, these updates will also include the parameters listed in Table 1. They will be named using the counter following the first decimal separator. The structure of the data product is designed to be flexible, allowing for the addition of different metadata fields and parameters. Further updates that include new parameters and/or require a new age modelling approach (i.e. no benthic $\delta^{18}\text{O}$ alignment) will be named using the counter before the first decimal separator. Any updates to add or correct (meta)data to an existing version that do not increase the number of sites will be indicated using a counter following the second decimal separator. Future versions will be made available on PANGAEA, and links to the updates will be provided at <https://doi.org/10.1594/PANGAEA.908831> (Jonkers et al., 2019).

8 Data availability

The PalMod 130k marine palaeoclimate data product can be downloaded in R, LiPD and netCDF format at <https://doi.org/10.1594/PANGAEA.908831> (Jonkers et al., 2019). The data can also be visualised and downloaded in LiPD and CSV formats at http://lipdverse.org/PalMod/current_version/. We encourage users of the data product to

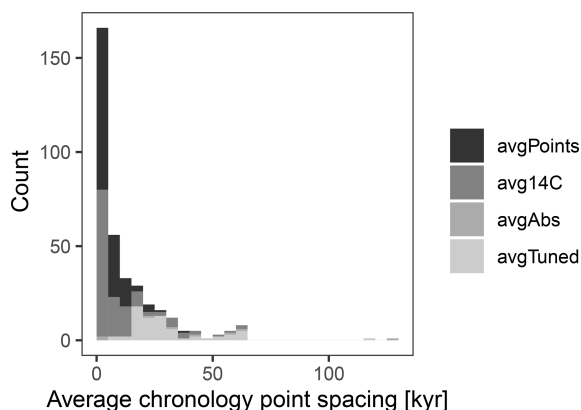


Figure 6. Average temporal spacing between chronology tie points (avgPoints). Tie points are also split into ^{14}C , absolute (tephra or palaeomagnetism) and tuned. The majority of the time series in the synthesis has a spacing of chronology tie points that is below 5000 years.

also cite the primary source of the data when using (individual time series of) this product.

9 Lessons learned: recommendations for data archiving

Data reuse and sharing are both made easier when data are archived in a standardised manner. Even though a large number of palaeoceanographic data are publicly available, metadata to facilitate interpretation of the raw or inferred palaeo-data are often not, or only partially, made available and need to be obtained from the original publication, which may be behind a paywall and not freely accessible. Synthesis efforts therefore still require a lot of time and effort to find, compile and standardise data and metadata. Only recently has the palaeoclimate community started to discuss data-archiving standards (Khider et al., 2019). However, implementation of the proposed Paleoclimate Community reporting Standard (PaCTS 1.0) will only affect new uploads to public repositories, and data already available (legacy data) are likely only going to be made compliant with the PaCTS through dedicated synthesis efforts. Below we list some of the main issues that we encountered during data synthesis. Our aims with mentioning these are to raise awareness of how the lack of standardisation affects data synthesis and thereby to encourage best practice in data reporting. We encourage researchers and also reviewers to treat data handling not as an afterthought but as an integral part of their study. After all, compared to generating the data, data handling is not a time-consuming task. Time spent on proper documenting and archiving is not wasted as it facilitates reuse of data and enables scientific progress in our field.

Disambiguate core names. An apparently trivial, but surprisingly common, first-order issue is that core names are

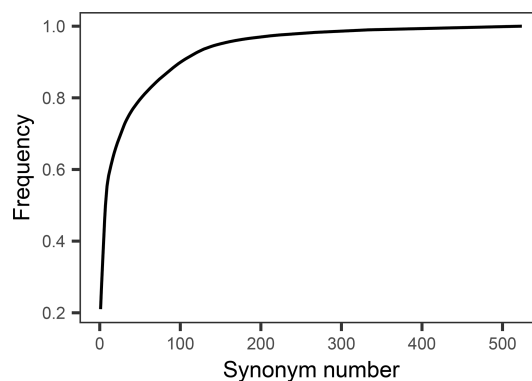


Figure 7. The problem of the absence of standardisation in parameter names. The cumulative frequency of synonyms for seawater temperature in our initial database (Sect. 2.2.1), showing that there are over 500 different names for the same parameter and that many of these are unique.

inconsistently archived. Different names for the same core arise not only from differences in hyphenation; truncation of (long) names; and minor variations in the same name that can, with expert knowledge, be linked but also from the use of altogether different names for the same core, e.g. reflecting differences in the labelling during an expedition and in the repository. This naming confusion renders it difficult to combine datasets from the same core, especially in an automated way, and to assess the uniqueness of time series from the same core for dereplication. We recommend using the full name as indicated in the cruise report where the core was first described.

Standardise vocabularies. Even though a vast number of palaeoclimate data are available in public repositories, the lack of standardisation of parameter vocabularies hinders efficient data processing. This problem is clearly illustrated by the fact that, for this synthesis, long synonym lists needed to be generated in order to group parameters. Each parameter in this synthesis had tens to hundreds of different names, of which many were unique (Fig. 7). This issue can be partly addressed by a consistent separation of parameter and attribute names (e.g. parameter $\delta^{18}\text{O}$, species *Globigerina bulloides*, instead of a single parameter “d18OGbul”), but even that calls for a standardisation of parameters and attributes.

Report sampling depth. All data reported here are based on measurements of discrete samples from a specific depth interval in a given core. Therefore, the synthesis requires information on the position of each sample. Since ages of the samples are always estimates and may differ among studies of the same archive, unambiguous information on sample depth is essential to reproduce and update the time series. Despite this, many studies fail to report sample depth and instead report only age. This problem is worse for spliced records that rely on a composite depth scale. Splicing approaches are often opaque, and original sample depths, or sample codes

that identify unique samples, are not always available. We recommend therefore that sample depth should be essential for palaeodata time series from (marine) sediments and that sample labels are archived for spliced records. This includes (I)OPD or DSDP sample labels (in full) or IGSN (if available).

Publish raw data. To ensure the reproducibility of inferred parameters (in this synthesis only temperature) and to ensure the harmonisation or updating of the calibration, the raw measured data are needed. Provided that the calibration is known, measured data can in some cases be calculated from the inferred data, but this is impossible for data based on microfossil transfer functions. Related to this is unclear information about the calibration that was used, particularly if the publication describing the calibration includes multiple different equations.

Include metadata. To assess ecological imprints on proxy signals (recording bias), temperature estimates as well as oxygen and carbon isotope data based on planktonic foraminifera also require that key metadata, such as species name, are archived in a standardised way. This is not universally carried out, and for example the species information is often only available in the original publication. Additional information to assess the uncertainty in proxy measurements, such as foraminifera shell size, the sample size (e.g. number of shells, concentration of alkenones) or reproducibility of repeat measurements, was often not available from the paper or from the archived data, and we encourage the archiving of such data in a standardised way. A similar issue applies to chronological data. Thanks to a longer history of reporting standards, radiocarbon (Stuiver and Polach, 1977) (meta)data are often rather complete. However, this information is often not included alongside the digitally available data, and for this synthesis a large proportion of the radiocarbon (meta)data had to be scraped from the literature. In our age modelling approach, we took reservoir age uncertainty into account, using data derived from the modelled reservoir ages (see Sect. 2.3). Alternative approaches are hindered by the fact that reservoir age uncertainty is almost never reported.

Avoid redundancy. A considerable amount of time was spent on the dereplication of time series of the same parameter from the same core that were archived multiple times. Repeat archiving happens when data are reused or, less commonly, updated. The dereplication task is not made easier by (incomplete or inconsistent) metadata reporting and can be avoided through better linking of existing datasets when, instead of re-uploading the data, the DOI or URL of the original data is provided.

Help rescue dark data. This synthesis is based on data that are publicly available, yet many palaeoceanographic time series are not archived in public repositories. Even though the proportion of this so-called dark data is by definition not known exactly, it likely affects every branch of palaeoceanography, and as a result palaeoceanographical data syn-

theses cannot be exhaustive. This problem is clearly exacerbated for syntheses relying on automated data mining. There are several shades of dark data, each requiring their own approach to retrieve them and make them available. Some data are only partially available, for instance datasets that lack sample depths. Such data only require additional data to make them reusable. These additional data can sometimes be calculated or obtained from cross-referencing different data files but in many cases will need to be retrieved from the data producers. Other datasets, in particular those from before the digital age, are presented in tables in the original publications. Progress has been made with digitising those datasets (especially in PANGAEA), but this work is not finished, and more effort is needed to make this data available to the community. There are also data that are used in publications but are not made available in any way (print or digital). This third shade of data can so far only be obtained from the original data producers or authors of the original publication or, if this proves impossible, needs to be digitised from graphs. Digitisation inevitably leads to a loss of accuracy of the data, and a dataset retrieved in this way should be flagged. A final category of dark data consists of data that are not part of a publication. Such datasets can be made publicly available and be associated with a DOI to ensure traceability. To reward data sharing, the use of data citations needs to be encouraged and data citations should be included in the evaluation of a researcher's impact.

Appendix A

Table A1. Data sources for version 1.0.0 of the PalMod 130k marine palaeoclimate data synthesis. An extended version of this table can be accessed at <https://doi.org/10.5281/zenodo.3739019> (Jonkers et al., 2020).

Site	Source
167_1017E	Kennett et al. (2000), Tada et al. (2000)
182_1132B	Holbourn et al. (2002), Brooks et al. (2002)
323_U1340A	Schlung et al. (2013)
90_594	Nelson et al. (1993), Wells and Okada (1997)
AHF_16832	Stott et al. (2000), Mortyn et al. (1996)
ASV13_1200	Duplessy et al. (2005), Ivanova (2002)
B997_328	Castañeda et al. (2004)
B997_330	Castañeda et al. (2004)
BP00_07_05	Simstich et al. (2008), Simstich et al. (2004)
CH69_K09	Govin et al. (2012), Cortijo et al. (1999), Labeyrie et al. (1999)
CHN82_24_4PC	Boyle and Keigwin (1985), Sarnthein et al. (1988), Ku et al. (1972)
EW9209_1JPC	Curry and Oppo (1997)
EW9209_2JPC	Curry et al. (1999)
EW9209_3JPC	Curry et al. (1999), Curry (1996)
EW9504_02PC	Stott et al. (2000)
EW9504_03PC	Stott et al. (2000)
EW9504_04PC	Stott et al. (2000)
EW9504_05PC	Stott et al. (2000)
EW9504_08PC	Stott et al. (2000)
EW9504_09PC	Stott et al. (2000)
FR01_97_12	Bostock et al. (2004), Bostock et al. (2009)
GeoB12615_4	Romahn et al. (2014)
GeoB13731_1	Fink et al. (2013)
GeoB1711_4	Little et al. (1997), Kirst et al. (1999), Vidal et al. (1999)
GeoB1720_2	Dickson et al. (2009)
GeoB3104_1	Arz et al. (1999), Arz et al. (1998)
GeoB3202_1	Arz et al. (1999)
GeoB3808_6	Jonkers et al. (2015)
GeoB4223_2	Freudenthal et al. (2002), Henderiks et al. (2002)
GeoB5844_2	Arz et al. (2003), Arz et al. (2007)
GeoB9508_5	Mulitza et al. (2008), Niedermeyer et al. (2009), Zarriess et al. (2011), Bouimetarhan et al. (2013)
GeoB9526_5	Zarriess and Mackensen (2011), Zarriess et al. (2011), Zarriess and Mackensen (2010)
GeoTu_SL148	Ehrmann et al. (2007)
GIK13289_2	Sarnthein et al. (1994)
GIK15612_2	Sarnthein et al. (1994), Kiefer (1998)
GIK15637_1	Sarnthein et al. (1994), Kiefer (1998), Zahn-Knoll (1986)
GIK17045_2	Sarnthein et al. (1994), Vogelsang et al. (2001)
GIK17045_3	Sarnthein et al. (1994)
GIK17049_6	Jung (1996), Vogelsang et al. (2001)
GIK17051_3	Jung (1996), Vogelsang et al. (2001)
GIK17940_2	Wang et al. (1999), Pelejero et al. (1999), Wang et al. (1999), Jian et al. (1999), Hu et al. (2012)
GIK17961_2	Wang et al. (1999), Pelejero et al. (1999)
GIK18471_1	Lo Giudice Cappelli et al. (2016)
GIK23258_2	Sarnthein et al. (2008), Martrat et al. (2003)
GIK23258_3	Sarnthein et al. (2008)
GIK23415_9	Jung (1996), Weinelt et al. (2003)
GIK23519_5	Millo et al. (2008)
GPC_5	Keigwin and Boyle (1999), Keigwin and Jones (1989), Keigwin and Jones (1994)
H214	Samson et al. (2005)
HLV02_02_17	Brunelle et al. (2007), Cook et al. (2005)
HU91_045_093	Hoogakker et al. (2014), Hoogakker et al. (2011)

Table A1. Continued.

Site	Source
KF13	Richter (1998)
KH94_3_LM_8	Oba and Murayama (2004)
KNR110_82	Sarnthein et al. (1988), Curry and Crowley (1987), Broecker et al. (1988)
KNR159_36	Oppo and Horowitz (2000), Curry and Oppo (2005), Carlson et al. (2008), Carlson et al. (2008)
KNR166_2_1	Lynch-Stieglitz et al. (2009)
KNR166_2_105	Lynch-Stieglitz et al. (2009)
KNR166_2_106	Lynch-Stieglitz et al. (2009)
KNR166_2_113	Lynch-Stieglitz et al. (2009)
KNR166_2_119	Lynch-Stieglitz et al. (2009)
KNR166_2_127	Lynch-Stieglitz et al. (2009)
KNR166_2_135	Lynch-Stieglitz et al. (2009)
KNR166_2_2	Lynch-Stieglitz et al. (2009)
KNR166_2_51	Lynch-Stieglitz et al. (2009)
KT90_9_21	Oba and Murayama (2004)
KT90_9_5	Oba and Murayama (2004)
LV28_42_4	Nürnberg and Tiedemann (2004), Kaiser (2001)
LV29_114_3	Riethdorf et al. (2013), Max et al. (2012), Max et al. (2014)
M35003_4	Rühlemann et al. (1999), Hüls (2000)
M77_2_056_5	Nürnberg et al. (2015), Mollier-Vogel et al. (2013)
M77_2_059_1	Nürnberg et al. (2015), Mollier-Vogel et al. (2013)
MD01_2378	Holbourn et al. (2005), Xu et al. (2006), Zuraida et al. (2009), Xu et al. (2008), Kawamura et al. (2006), Dürkop et al. (2008), Sarnthein et al. (2011)
MD01_2416	Sarnthein et al. (2004), Gebhardt et al. (2008), Gray et al. (2018a), Sarnthein et al. (2015), Sarnthein et al. (2013)
MD02_2489	Gebhardt et al. (2008)
MD02_2575	Ziegler et al. (2008), Nürnberg et al. (2008)
MD02_2589	Diz et al. (2007), Molyneux et al. (2007)
MD06_2986	Ronge et al. (2015)
MD06_2990	Ronge et al. (2015)
MD06_3067	Bolliet et al. (2011)
MD06_3075	Fraser et al. (2014)
MD07_3076	Gottschalk et al. (2015), Gottschalk et al. (2016), Waelbroeck et al. (2011), Roberts et al. (2016), Skinner et al. (2010)
MD73_025	Cline et al. (1984), Labeyrie and Duplessy (1985), Labracherie et al. (1989)
MD84_527	Sarnthein et al. (1988), Pichon et al. (1992), Duplessy et al. (1988), Labracherie et al. (1989)
MD88_770	Labeyrie et al. (1996)
MD95_2010	Dokken and Jansen (1999), Risebrobakken et al. (2005), Govin et al. (2012)
MD95_2024	Hoogakker et al. (2011), Hoogakker et al. (2014), Weber et al. (2001), Korte and Hesselbo (2011)
MD95_2039	Thomson et al. (1999), Salgueiro et al. (2014), Eynaud et al. (2009), Schönfeld et al. (2003)
MD95_2040	Voelker and de Abreu (2011), de Abreu et al. (2003), Pailler and Bard (2002), Voelker et al. (2009), Moreno et al. (2002), Schönfeld et al. (2003)
MD95_2043	Cacho et al. (2006), Cacho et al. (1999), Martrat et al. (2014)
MD96_2098	Pichevin et al. (2005), Daniau et al. (2013)
MD97_2120	Sachs and Anderson (2005), Pahnke and Sachs (2006), Pahnke (2003), Pahnke (2005)
MD98_2181	Stott et al. (2007), Saikku et al. (2009), Stott (2002), Stott (2007)
MD99_2236	Jennings et al. (2015)
MD99_2339	Voelker et al. (2006), Voelker et al. (2009), Voelker and de Abreu (2011)
MD99_2343	Sierro et al. (2005), Frigola et al. (2008)
MSM05_5_712_1	Werner et al. (2011), Spielhagen et al. (2011)
MSM05_5_712_2	Werner et al. (2013), Müller et al. (2012), Müller and Stein (2014)
MSM05_5_723_2	Werner et al. (2016), Müller et al. (2012)
MV0502_4JC	Waddell et al. (2009)
NA87_22	Duplessy et al. (1992), Vogelsang et al. (2001), Gherardi et al. (2009)
NEAP_04K	Rickaby and Elderfield (2005), Hall et al. (2004)
ODP1145	Oppo and Sun (2005)

Table A1. Continued.

Site	Source
ODP846	Mix et al. (1995), Lawrence (2006), Martinez et al. (2003)
ODP980	Oppo et al. (2003), McManus et al. (1999), Oppo et al. (2006), Ortiz et al. (1999)
ODP984C	Praetorius et al. (2008), Came et al. (2007), Ortiz et al. (1999)
P7	Pedersen et al. (1988), Pedersen et al. (1991)
POS200_10_6_2	Baas et al. (1997)
PS1730_2	Nam (1997)
PS1878_3	Telesinski et al. (2014), Telesinski et al. (2013)
PS2138_1	Wollenburg et al. (2001), Knies and Vogt (2003), Knies and Stein (1998), Nowaczyk et al. (2003), Knies et al. (2000)
PS75_059_2	Ullermann et al. (2016), Lamy et al. (2014), Ronge et al. (2016)
R657	Weaver et al. (1998), Sikes et al. (2002)
RAPiD_15_4P	Thornalley et al. (2010), Thornalley et al. (2010), Thornalley et al. (2011)
RAPiD_17_5P	Thornalley et al. (2010), Thornalley et al. (2011)
RC11_83	Charles et al. (1996), Charles and Fairbanks (1992), Piotrowski et al. (2004)
RC16_119	Oppo and Horowitz (2000)
RC16_84	Oppo and Horowitz (2000)
RS147_GC07	Sikes et al. (2009)
SK157_14	Ahmad et al. (2008)
SO136_003GC	Ronge et al. (2015), Barrows et al. (2007)
SO164_17_2	Bahr et al. (2011)
SO201_2_12KL	Riethdorf et al. (2013), Max et al. (2012), Max et al. (2014)
SO201_2_85	Riethdorf et al. (2013), Riethdorf et al. (2013), Max et al. (2014), Max et al. (2012), Max et al. (2014), Riethdorf et al. (2016)
SO213_2_59_2	Tapia et al. (2015)
SO213_2_82_1	Ronge et al. (2015), Ronge et al. (2016)
SO213_2_84_1	Ronge et al. (2015), Ronge et al. (2016)
SO42_74KL	Sirocko (2000), Sirocko et al. (1993), Sirocko et al. (1991), Kim et al. (2004), Schulz (1995)
SO82_5_2	Jung (1996), van Kreveland et al. (2000)
SU81_18	Bard et al. (1989), Bard (2000), Waelbroeck et al. (2001)
SU90_11	Labeyrie et al. (1995), Jullien et al. (2006)
SU90_24	Elliot et al. (2002), Elliot et al. (1998), Elliot et al. (2001)
TR163_22	Lea et al. (2006)
V19_27	Lyle et al. (2002), Koutavas and Lynch-Stieglitz (2003)
V19_28	Lyle et al. (2002), Koutavas and Lynch-Stieglitz (2003)
V19_30	Shackleton and Pisias (2013), Lyle et al. (2002), Bond (1997, 1976)
V23_81	Jansen and Veum (1990), Broecker et al. (1988), CLIMAP project members (1981)
V24_253	Oppo and Horowitz (2000)
V25_59	Sarnthein et al. (1988), Waelbroeck et al. (1998), CLIMAP project members (1981)
V28_14	CLIMAP project members (1981), Kellogg et al. (1978)
V29_202	Oppo and Lehman (1995)
W8709A_13	Lyle et al. (1992), Lyle et al. (2000), Kienast et al. (2002), Gardner et al. (1997), Lund and Mix (1998)
WIND_28K	McCave et al. (2005), Johnstone et al. (2014), Kiefer et al. (2006)
Y69_106	Lyle et al. (2002)

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